

Forcing variables in simulation of transpiration of water stressed plants determined by principal component analysis**

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Abstract. To date, measuring plant transpiration at canopy scale is laborious and its estimation by numerical modelling can be used to assess high time frequency data. When using the model by Jacobs (1994) to simulate transpiration of water stressed plants it needs to be reparametrized. We compare the importance of model variables affecting simulated transpiration of water stressed plants. A systematic literature review was performed to recover existing parameterizations to be tested in the model. Data from a field experiment with common bean under full and deficit irrigation were used to correlate estimations to forcing variables applying principal component analysis. New parameterizations resulted in a moderate reduction of prediction errors and in an increase in model performance. Ag_s model was sensitive to changes in the mesophyll conductance and leaf angle distribution parameterizations, allowing model improvement. Simulated transpiration could be separated in temporal components. Daily, afternoon depression and long-term components for the fully irrigated treatment were more related to atmospheric forcing variables (specific humidity deficit between stomata and air, relative air humidity and canopy temperature). Daily and afternoon depression components for the deficit-irrigated treatment were related to both atmospheric and soil dryness, and long-term component was related to soil dryness.

Key words: Ag_s model, net CO_2 assimilation rate, stomatal conductance, common bean

INTRODUCTION

Measuring plant transpiration rate at canopy scale is laborious and results are often uncertain. Modern techniques to measure plant transpiration include porometers

(Ansley *et al.*, 1994; Silva *et al.*, 2016), gas analysers (Escalona *et al.*, 2000) and heat balance sap flow gauges (De Lorenzi and Rana, 2000). While the last one is more suitable for trees and difficult to apply in annual crops, the porometers and gas analyser systems take measurement from the leaf itself (Dugas *et al.*, 1993). However, the *in situ* use of these instruments is limited because obtaining measurements during the entire crop cycle and at a high time frequency is unfeasible. To tackle the problem, estimates of plant transpiration rates by numerical modelling can be used as a tool to assess high time frequency plant transpiration data (Oliosio *et al.*, 2005). In these models, input variables such as air or canopy temperature are used that can be measured easier and at higher frequency than transpiration.

A physical-mechanistic way to model the plant transpiration rate is to account for the conductance (or its inverse: resistance) of the water dissipation pathway through the canopy to the atmosphere and to relate the conductance to driving environmental forces like evaporation and total canopy leaf area. The leaf stomatal conductance is the most variable parameter on the pathway and in a physical approach of transpiration modelling it is related to the stomatal conductance, allowing to determine the inherent parameters of the process. This sort of relation was reported by pioneer physiologists like Farquhar *et al.* (1980) and Goudriaan *et al.* (1985), and derives from the observation that the exchange of water between the leaves and

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the atmosphere takes place in the gaseous phase of water, mostly through the stomata. Both in experimental studies (Cowan, 1982) and by numerical modelling (Jacobs *et al.*, 1996), stomatal conductance has been shown to be related to net CO₂ assimilation rate A , to the water vapour pressure deficit in the atmospheric air, to the specific vapour deficit between the intercellular leaf air spaces and the atmospheric air D_s , and to the intercellular CO₂ concentration C_i . In Ag_s models (A being the net CO₂ assimilation rate and g_s the stomatal conductance to water vapour), the stomatal conductance to water vapour is derived based on the net CO₂ assimilation rate calculated from mainly physically based parameterizations of the above variables (Jacobs, 1994; Albergel *et al.*, 2010).

The Ag_s model proposed by Jacobs (1994) is based on the equations by Goudriaan *et al.* (1985), and describes the stomatal processes by parameterising the net CO₂ assimilation responses to the environmental factors such as air humidity and temperature and radiation. However, as each of the variables affects the sensitivity of stomata to other factors, the model also describes the synergistic interactions between different stimuli. The model by Jacobs (1994) has been widely used in numerical schemes accounting for photosynthesis and carbon fluxes due mainly to its physical plant-environment representation. Examples are the meteorological model ECMWF (Boussetta *et al.*, 2013) or the crop model WOFOST-SWAP (Supit *et al.*, 2012; Van Dam *et al.*, 2008), but its possible application in plant transpiration simulations is still open.

Jacobs (1994) used parameterizations disregarding the effect of water stress in plants by high atmospheric demand of water vapour or by soil water shortage. However, studies show that water stress affects the stomatal conductance, the CO₂ assimilation and, consequently, introduces limits to the productivity and growth of plants (Lipiec *et al.*, 2013). According to Galle *et al.* (2009), the limiting photosynthesis factors under these conditions and their possible interactions with other environmental conditions are still not very well understood. The increase in stomatal resistance of plant leaves to water vapour induced by water stress has been considered as equally limiting for CO₂ assimilation, since the diffusion of CO₂ from the carboxylation sites in the chloroplast to the atmosphere is hampered. However, the reduced CO₂ assimilation rates during water stress events can be explained also by other factors, like internal leaf restrictions in the CO₂ path from intercellular spaces to the mesophyll cells, to the chloroplasts and to the carboxylation sites (Galle *et al.*, 2009).

The contribution of each internal leaf compartment to the reduction of CO₂ assimilation during water stress events is not completely quantified (Warren, 2008). Observational studies show that both the CO₂ compensation point Γ , mesophyll conductance g_m , maximum CO₂ assimilation rate $A_{m,max}$ and dark respiration of photosynthesis R_d vary significantly depending on the plant water status.

One of the first studies dealing with the effects of water stress on the CO₂ compensation point Γ made by Smolander and Lappi (1984) showed Γ to increase with the intensification of water stress and increasing leaf temperature in willow trees. Recently, a review by Srikanta Dani *et al.* (2015) indicated a general increase of leaf temperature and Γ as a response of stomatal closure due to water stress. In sunflower plants, Tezara *et al.* (1999) observed $A_{m,max}$ to decrease with water stress intensification whereas Γ progressively increased over the same period. In vineyards, Escalona *et al.* (1999) reported a pronounced effect of soil water shortage on $A_{m,max}$ with a reduction of 60% in relation to well watered plants. A decrease of $A_{m,max}$ of the drought-sensitive European beech (*Fagus sylvatica*) tree species was also reported by Hommel *et al.* (2016). The mesophyll conductance g_m decreased in response to water stress and to the abscisic acid production in studies by Flexas *et al.* (2006) and Galle *et al.* (2009), but g_m can also change solely in response to water stress, usually decreasing as observed by Warren (2008). These results corroborate the hypothesis that g_m plays an important role in the photosynthetic response of plants to climatic forcing (Flexas *et al.*, 2008). The effect of water stress on the dark respiration R_d is inconclusive. Some studies show a decrease of R_d as a function of water stress (Galmés *et al.*, 2007), whereas others report an increase (Zagdańska, 1995). According to Ribas-Carbo *et al.* (2005), changes in R_d are smaller than in photosynthesis, causing a significant increase in the ratio of respiration/photosynthesis under water stress, and indicating the role of respiration to become more important with increasing water stress.

In the Ag_s model by Jacobs (1994), the variables Γ , g_m , $A_{m,max}$ and R_d , fundamental to calculate the net CO₂ assimilation, are empirically treated as a function of leaf temperature, limiting the model representation of plant physiology during the occurrence of water stress. To estimate the transpiration rate of a crop canopy, the net CO₂ assimilation for each vertical canopy layer (leaf scale) is calculated by the extinction of photosynthetically active radiation I_{PAR} inside the canopy, since the leaves at the top of the canopy intercept the majority of the downward solar radiation, reducing the photosynthetic activity of the lower canopy layers. The I_{PAR} extinction inside the canopy follows Beer law of exponential extinction considering the angular distribution of leaves, as proposed by Roujean (1996). Under conditions of water stress, however, the extinction of I_{PAR} by plants is no longer exponential because morphologic aspects of the leaves, like the angular distribution G , are modified (Archontoulis *et al.*, 2011) and the Ag_s model by Jacobs (1994) is no longer adequate.

In this context, it was the objective of this study to shed more light on the relative importance of the variables affecting transpiration simulation of plants exposed to water stress. To do so, we compare modified versions of the Jacobs (1994) model with specific parameterizations of

Γ , g_m , $A_{m,max}$, R_d and G for dry conditions to the original Jacobs (1994) model. We use data from a field experiment in Brazil and correlate estimated plant transpiration rates to relevant environmental and plant variables by a principal component analysis.

MATERIAL AND METHODS

Experimental data were obtained from a 1000 m² field plot, cropped with common bean (*Phaseolus vulgaris* L., Carioca cultivar), in Piracicaba, São Paulo State, Brazil (22° 42' S, 47° 38' E) between June and September, 2010. The soil at the experimental plot is a Rhodic Kanhapudalf, according to USDA Soil Taxonomy. Bulk density is 1560 kg m⁻³ in the Ap horizon (0-0.2 m) and 1380 kg m⁻³ in the Bt horizon (0.2-0.8 m). Clay content is 0.45 and 0.55 kg kg⁻¹ in Ap and Bt horizon, respectively. Local climate is classified as humid subtropical, with dry winters and hot summers (Köppen Cwa), favourable for the cultivation of Common Bean. The annual rainfall is approximately 1300 mm. However, in the dry winter months (June – August), rainfall is almost absent and irrigation is required to successfully grow crops.

The area was divided in two subplots: one irrigated by sprinklers during the entire crop cycle (fully irrigated treatment – FI), the other one subject to water stress in the reproductive phase between August 4 and September 1 (deficit irrigated treatment – DI). The time interval between August 4 and September 2 was defined as the study period in this paper as it was the period during the crop cycle when plants on the two plots were subjected to distinctly different soil water conditions. Relative air humidity (H_p , %) and air temperature (t_{air} , °C) were measured at a height of 2 m above the centre of each irrigation treatment, using a Campbell Scientific CS215L9 probe connected to a CR1000 data-logger. Canopy temperature (t_c , °C) was measured via two automated infrared thermometers (Apogee model SI111, target area of 65 m²), one per treatment, connected to the same data-logger. For measurements of soil water pressure head (h , MPa), polymer tensiometers (Bakker *et al.*, 2007; Van der Ploeg *et al.*, 2008) were used, with a measurement range between 0 and 1.47 MPa (-150 m). Soil water pressure head was measured at 0.05, 0.1 and 0.3 m depth at two observation points per subplot (two observation points (identified as 1 and 2) in the fully irrigated treatment, and two observation points (3 and 4) in the deficit irrigated treatment). All measurements were performed every 30 min averaging the previous period. To avoid including the impact of irrigation water on temperatures and atmospheric vapour pressure in the data set, data obtained during irrigation events were excluded from further analysis.

The leaf area index (L_{LAI} , m² m⁻²) was measured by an indirect and non-destructive method with a ceptometer. Plant Canopy Analyser, model LAI2000® from Li-Cor. LAI measurements were taken five times per each irriga-

tion treatment during the experimental period, on August 4, 12, 19 and 27 and on September 2. Daily values of L_{LAI} were obtained by linear interpolation between the measurements, and ranged from 3.37 to 5.70 m² m⁻² in the fully irrigated treatment and between 2.49 and 3.37 m² m⁻² in the deficit irrigated treatment.

During the 90 days of the field experiment (June 15 – September 13), rainfall was observed on July 13, 14 and 15 (63 mm in total) and then again, after more than 50 days, on September 7 (13 mm) during the ripening period. In the periods between August 2 and 23 and between August 25 and September 1, only the fully irrigated treatment was irrigated. On August 24, a small irrigation (~ 15 mm) was performed on the deficit irrigated treatment to guarantee the survival of the crop. In total (rainfall + irrigation), the fully irrigated treatment received 426.5 mm of water and the deficit irrigated treatment received 314.5 mm. Additional measurements and experimental details are described in Durigon *et al.* (2012).

Plant transpiration simulations were performed using the Ag_s model by Jacobs (1994) and Jacobs *et al.* (1996). The theoretical description of model parameterizations is presented in the original publications. The Ag_s model was numerically programmed using Fortran programming language and the source code was described in Durigon (2011). In the original model, the relationship between stomatal conductance to water vapour g_s , expressed in mm s⁻¹, and net CO₂ assimilation A (mg m² s⁻¹) is given, as a first approximation, by:

$$g_s = 1.6 g_{sc} = 1.6 10^3 \frac{A}{C_s - C_i}, \quad (1)$$

where: the factor 1.6 is the ratio between air diffusivities of H₂O and CO₂, g_{sc} (mm s⁻¹) is the stomatal conductance to CO₂, C_s and C_i (mg m⁻³) are the CO₂ concentrations at the leaf surface and the intercellular air spaces in the leaves, respectively, parameterised as a function of specific humidity deficit between leaf and air, D_s (mg kg⁻¹), defined as:

$$D_s = q_s(t_c) - q_{air}. \quad (2)$$

In this equation, $q_s(t_c)$ is the specific humidity at saturation (mg kg⁻¹) as a function of canopy temperature t_c and q_{air} is the specific humidity of the atmospheric air near leaves (mg kg⁻¹).

Assuming a vertical distribution of leaves, net CO₂ assimilation A and stomatal conductance g_s are given by:

$$A = \frac{L_{LAI}}{h_p} \int_0^h A dz, \quad (3)$$

$$g_s = \frac{L_{LAI}}{h_p} \int_0^h g_s dz,$$

in which L_{LAI} (m² m⁻²) is the leaf area index, h_p (m) is

plant height and dz (m) is a height interval. The integrals are solved by applying a five-point Gaussian quadrature scheme *eg*:

$$A = \frac{L_{LAI}}{h_p} \sum_{i=1}^5 W_i A(z_i),$$

$$g_s = \frac{L_{LAI}}{h_p} \sum_{i=1}^5 W_i g_s(z_i),$$
(4)

where: z_i and W_i are distance and weight of point i , respectively, and values are presented in Durigon (2011).

The plant transpiration rate T ($\text{mg m}^{-2} \text{s}^{-1}$) is determined by the total conductance to the water vapour flux g_t (mm s^{-1}):

$$T = 10^{-3} \rho_{air} g_t D_s L_{LAI},$$
(5)

$$g_t = g_s + g_m + 2g_{bl},$$
(6)

where: ρ_{air} is the air density (1.2 kg m^{-3} at 15°C) and g_{bl} (mm s^{-1}) is the boundary layer conductance around leaf calculated by:

$$g_{bl} = k \left(\frac{u}{W_l} \right)^{0.5},$$
(7)

in which k is an empirical constant ($5.6 \text{ mm s}^{-0.5}$), u is the wind speed (mm s^{-1}) and W_l is the leaf width parallel to wind (assumed to be 100 mm). The cuticular conductance was considered negligible when compared to other conductances and was not included in Eq. (6). Multiplication of g_{bl} by 2 allows accounting for both sides of leaves.

As mentioned in the introduction, the variables Γ , g_m , $A_{m,max}$, R_d and G are treated in a semi-empirical way in the Ag_s model by functions shown in Fig. 1 as solid black lines. The response of Γ to the temperature is given by a Q_{10} function, a simplification of the Arrhenius function:

$$\Gamma = \Gamma_{25} Q_{10}^{\frac{t_c - 25}{10}},$$
(8)

in which t_c is the canopy temperature in $^\circ\text{C}$, and Γ_{25} is the value of Γ at $t_c = 25^\circ\text{C}$. Both Γ_{25} and Q_{10} for Γ estimation are presented in Table 1.

To obtain g_m and $A_{m,max}$, a generic equation is used:

$$X(t_c) = \frac{x_{25} Q_{10}^{\frac{t_c - 20}{10}}}{(1 + \exp(0.3(t_1 - t_c)))(1 + \exp(0.3(t_c - t_2)))},$$
(9)

where: $X(t_c)$ is the value of g_m or $A_{m,max}$ at temperature t_c , t_1 and t_2 represent reference temperatures and need to be adjusted to minimise the specific species characteristics (*eg* the optimum temperature for plants with a C_3 photosynthetic pathway is lower than for plants with a C_4 photosynthetic pathway), and X_{25} is the value of g_m or $A_{m,max}$ at $t_c = 25^\circ\text{C}$ ($g_{m,25}$ and $A_{m,max,25}$). Both Q_{10} , X_{25} and reference temperatures to the g_m and $A_{m,max}$ estimation by the original Jacobs model are presented in Table 1.

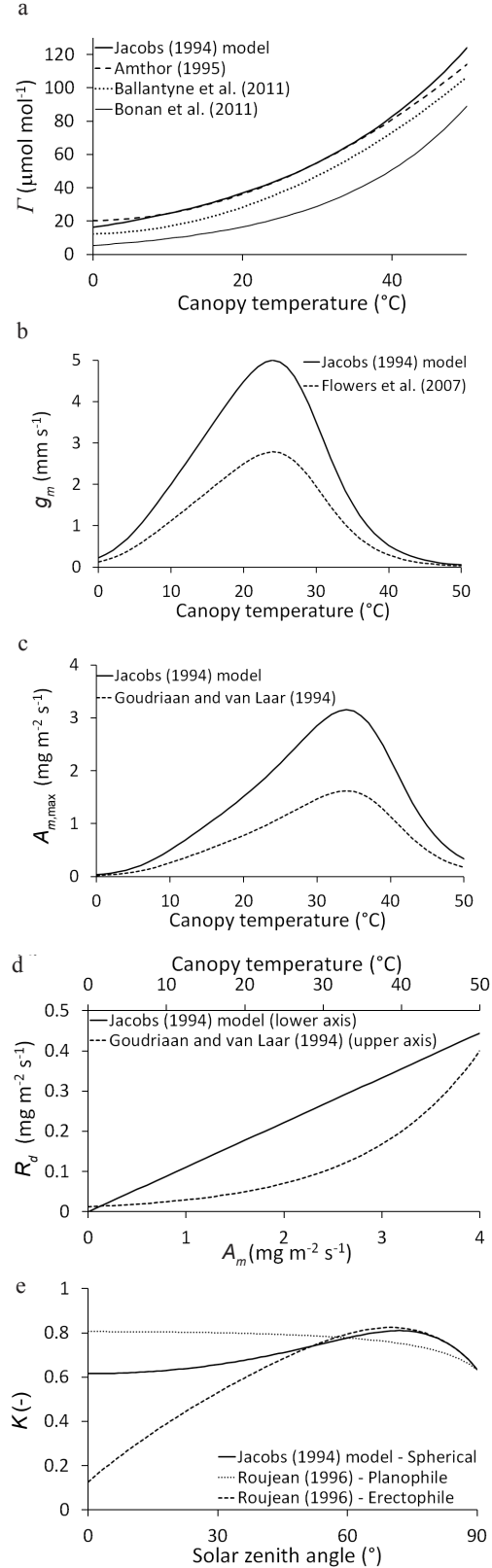


Fig. 1. Parameters used by the original Ag_s model by Jacobs (1994) and in this study (Γ – CO_2 compensation point, g_m – mesophyll conductance, $A_{m,max}$ – maximum CO_2 assimilation rate, R_d – dark respiration of photosynthesis, K – total light extinction coefficient (function of leaf angle distribution G)).

Table 1. Parameters used by the original Ag_s model by Jacobs (1994) and used in this study

Parameter (X)	X_{25}	Q_{10}	t_1 (°C)	t_2 (°C)
Γ ($\mu\text{mol mol}^{-1}$)	45	1.5	–	–
$A_{m,\text{max}}$ ($\text{mg m}^{-2} \text{s}^{-1}$)	2.2	2	8	38
g_m (mm s^{-1})	7	2	5	28

To estimate R_d , the following relation is used:

$$R_d = \frac{A_m}{9}, \quad (10)$$

where: A_m ($\text{mg m}^{-2} \text{s}^{-1}$) is the photosynthetic rate at saturating light intensity.

To upscale the Ag_s model by Jacobs (1994) from leaf to canopy transpiration, a spherical leaf angle distribution was considered, represented by parameter G ($G = 0.5$) in the extinction coefficient of direct light $K_{dr}(z)$ (Calvet *et al.*, 1998):

$$K_{dr}(z) = 1 - \exp\left(-\frac{G}{\cos(\theta_s)} b L_{LAI} \frac{h_p - z}{h_p}\right), \quad (11)$$

in which θ_s (°) is the solar zenith angle, b is the foliage scattering coefficient, h_p (m) is the plant height and z (m) is a level above the soil surface. The total light extinction coefficient $K(z)$ between the top of the canopy and the level z is expressed by:

$$K(z) = f(\theta_s) K_{df}(z) + (1 - f(\theta_s)) K_{dr}(z), \quad (12)$$

where: $K_{df}(z)$ is the extinction coefficient of diffuse light and $f(\theta_s)$ is the ratio of diffuse to total solar radiation at the top of the canopy which is given by:

$$f(\theta_s) = \frac{0.25}{0.25 + \cos(\theta_s)}. \quad (13)$$

A systematic literature review was performed to retrieve parameterizations or parameter values of Γ , g_m , $A_{m,\text{max}}$, R_d and G different from the ones used in the original Ag_s

model and which could better represent the environmental water stress condition. A systematic literature review is a standard procedure performed for the survey of available information (Ganann *et al.*, 2010). Our systematic literature review consisted of the combination of the terms ‘CO₂ compensation point’, ‘mesophyll conductance’, ‘maximum CO₂ assimilation’, ‘dark respiration’, and ‘leaf angle distribution’ to each of the terms ‘parameterization’ and ‘C₃’ (referring to C₃ plants), and a main search was performed in the Google Scholar database of June 2015, defining the specific publication time interval from 1990 to 2015. The number of publications retrieved for each combination of terms is listed in Table 2. Among these publications, a detailed selection was performed to identify parameterizations that could be used in the Ag_s model. Compiling these publications, a final set of 8 parameterizations was obtained and added to the Ag_s model. The number of publications selected and the number of parameterizations added to the Ag_s model for each combination of terms are listed in Table 2. Although a relatively high number of publications presented parameterizations that could be used in the model (107 publications in total), most of these refer to the same parameterizations. This indicates that a few mathematical representations of the biochemical variables related to the CO₂ assimilation process are actually used in crop modelling.

The functions representing Γ , g_m , $A_{m,\text{max}}$, R_d and G in the original Jacobs (1994) model were replaced by functions obtained for C₃ plants in the systematic literature review. The new functions added to the model are also shown in Fig. 1. To modify the Γ function in the original Ag_s model, three new functions were used (Fig. 1a). In the first one, described by Amthor (1995), the CO₂ compensation point Γ is a quadratic function of canopy temperature t_c (°C):

$$\Gamma = 44.7 + 1.88(t_c - 25) + 0.036(t_c - 25)^2. \quad (14)$$

A second, very similar function was proposed by Ballantyne *et al.* (2011):

$$\Gamma = 36.9 + 1.88(t_c - 25) + 0.036(t_c - 25)^2. \quad (15)$$

Table 2. Number of publications retrieved in the systematic literature review, publications selected and parameterizations added to the Ag_s model for combination of the search term combined to terms ‘parameterization’ and ‘C₃’

Term combined to ‘parameterization’ and ‘C ₃ ’ terms	Publications retrieved	Publications selected	Parameterizations added to the Ag_s model
CO ₂ compensation point	196	51	3
Mesophyll conductance	190	17	1
Maximum CO ₂ assimilation	10	7	1
Dark respiration	493	17	1
Leaf angle distribution	171	16	2

The third function was presented by Bonan *et al.* (2011) who estimated Γ from the dependence of the Michaelis-Menten coefficients to O₂ and CO₂ diffusion to the canopy temperature:

$$\Gamma = 0.5 \frac{K_c}{K_o} 0.21O, \quad (16)$$

$$K_c = K_{c25} f(t_c); \quad Q_{10} = 2.1, \quad (17)$$

$$K_o = K_{o25} f(t_c); \quad Q_{10} = 1.2, \quad (18)$$

$$f(t_c) = Q_{10}^{\frac{(t_c - 298.15)}{10}}, \quad (19)$$

where: K_c and K_o are the Michaelis-Menten coefficients for CO₂ and O₂, respectively, O is the atmospheric O₂ concentration (0.209 mol mol⁻¹), K_{c25} (30 Pa) and K_{o25} (30 000 Pa) are the Michaelis-Menten coefficients for CO₂ and O₂ at 25°C, respectively, $f(t_c)$ is the temperature dependence function of K_c and K_o , and t_c is the absolute canopy temperature (K). In the normal canopy temperature range, all three functions represent a similar increase of Γ , but slightly greater Γ values are obtained from the functions of the original model, of Amthor (1995) and of Ballantyne *et al.* (2011), as compared to the Bonan *et al.* (2011) function. All of them are supported by observations presented by Smolander and Lappi (1984) who showed Γ to increase with the intensification of water stress and increasing leaf temperature in willow trees (*Salix babylonica*).

The original Jacobs (1994) simulations are based on $g_m = 7$ mm s⁻¹ under ideal soil hydraulic conditions at 25°C, in agreement with Nobel (1991). Flowers *et al.* (2007) reported a value of $g_m = 3.9$ mm s⁻¹ for C₃ plants under water stress at 25°C. We replaced the original value of 7 mm s⁻¹ by 3.9 mm s⁻¹, which introduced a large change in the g_m function (Fig. 1b). The maximum value of g_m equals 5 mm s⁻¹ in the original model, but using the Flowers *et al.* (2007) observation it reduces to 2.79 mm s⁻¹, both at $t_c = 24$ °C.

For $A_{m,max}$, the original function was replaced by:

$$A_{m,max}(t_c) = \frac{A_{m,max20} Q_{10}^{\frac{t_c - 20}{10}}}{(1 + \exp(0.3(t_1 - t_c)))(1 + \exp(0.3(t_c - t_2)))}, \quad (20)$$

in which $A_{m,max}$ at 20°C ($A_{m,max20}$) is assumed equal to 0.8 mg m⁻² s⁻¹, replacing the value for 25°C, and the exponential of the Q_{10} function was modified following Goudriaan and van Laar (1994) (Fig. 1c). Using the new parameterization, the maximum $A_{m,max}$ is equal to 1.62 mg m⁻² s⁻¹ while the maximum $A_{m,max}$ in the original model is equal to 3.15 mg m⁻² s⁻¹, both at $t_c = 34$ °C. The lower maximum of the new $A_{m,max}$ function curve could indicate that it better represents the water stress condition following observations of Tezara *et al.* (1999) and Escalona *et al.* (1999), who reported $A_{m,max}$ to decrease in water stress occurrences.

R_d as a function of canopy temperature was estimated using the function proposed by Goudriaan and van Laar (1994):

$$R_d = R_{d20} Q_{10}^{\left(\frac{t_c - 20}{10}\right)}, \quad (21)$$

where: R_{d20} is dark respiration at 20°C (= 0.05 mg m⁻² s⁻¹) and Q_{10} was taken as 2 (Fig. 1d). While the original R_d parameterization is a linear function of A_m , the modified function is exponential with t_c , but both parameterizations agree with the observations of Zagdańska (1995) who indicated R_d to increase as plants became water stressed, closing stomata and reducing transpiration, causing higher leaf temperatures.

Leaf angle distribution was modified from the spherical, as used in Jacobs (1994) distributions suggested by Roujean (1996), to planophile:

$$G = \cos\theta_s, \quad (22)$$

or to erectophile:

$$G = \frac{2}{\pi} \sin\theta_s. \quad (23)$$

The three distributions are illustrated in Fig. 1e. The spherical function implies lower sensitivity to the solar zenith angle q_s of light absorption by the canopy (related to the light extinction coefficient K). Using $b = 0.9442$, $L_{LAI} = 4$ m² m⁻², $h_p = 0.6$ m and $z = 0.4$ m, the resulting K is around 0.7. The planophile function simulates higher light interception, making it less sensitive to the solar zenith angle. The erectophile function, on the other hand, reduces interception, making it more sensitive to the solar zenith angle, extinguishing less light as θ_s decreases. This is, in fact, a behaviour adopted by plants under water stress to avoid an excessive radiation load during the warmest hours of the day. Atti *et al.* (2005), for example, experimentally observed that soybean plants changed their leaf angle distribution of spherical to erectophile as a function of soil water shortage.

To evaluate the performance of the simulations, predicted canopy temperatures were compared to measurements. Canopy temperature was calculated by the energy balance approach, using hourly net radiation measured in the weather station near the experiment, the transpiration rate estimated by the A_g model to calculate the latent heat flux, and assuming the soil heat flux to be equal to zero in fully closed vegetated surfaces (Allen *et al.*, 1998).

Three statistical indices were used for model performance quantification:

– root-mean-square error of prediction ($RMSEP$, °C):

$$RMSEP = \sqrt{\frac{\sum_{i=1}^n (t_{c,i} - \hat{t}_{c,i})^2}{n}}, \quad (24)$$

– mean error (ME , °C):

$$ME = \frac{1}{n} \sqrt{\sum_{i=1}^n (t_{c,i} - \hat{t}_{c,i})^2}, \quad (25)$$

– efficiency coefficient E (Nash and Sutcliffe, 1970):

$$E = 1 - \frac{\sum_{i=1}^n (t_{c,i} - \hat{t}_{c,i})^2}{\sum_{i=1}^n (t_{c,i} - \bar{t}_{c,i})^2}, \quad (26)$$

where: t_c and \hat{t}_c are observed and estimated values of canopy temperature, respectively, \bar{t}_c is the mean of observed values and n is the total number of observations.

Many environmental (soil and atmosphere) and plant variables act together to determine the plant transpiration rate. The relative importance of each of these components can be evaluated using a principal component analysis (PCA) (Abdi and Williams, 2010), a specific statistical methodology to explain a process governed by a large set of variables in terms of a smaller set. It allows the interpretation of a process by comprising the maximum information of an initial data set measured through the total variance of two or three representative components with an appropriate function of the total variance. The add-in NumXL available in Microsoft Excel 2013 was used to perform the PCA.

The principal component analysis (PCA) was performed in two ways for both irrigation treatments: by taking the transpiration rate data series together with the environmental and plant input variables, and by taking only the data series of the environmental and plant input variables (named as EP), aiming to characterise the plant transpiration controlling environment. The environmental and plant variables used in the PCA were the specific humidity deficit between stomata and air (D_s), the photosynthetically active radiation (I_{PAR}), the relative humidity (H_R), canopy and air temperature (t_c and t_{air}), the wind speed (u), and the leaf area index (L_{LAI}). In addition, time series of the atmospheric water potential (ψ , MPa) and the mean soil water pressure head (h_m , MPa) were included in the PCA as environmental variables. The atmospheric water potential and the soil water pressure head decrease when air and soil get dry, and the difference between these two variables, represent the main driving force to water flow in the soil-plant-atmosphere continuum.

The atmospheric water potential ψ (MPa) can be calculated from air absolute temperature t_{AIR} (K) and relative humidity H_R (here in Pa Pa⁻¹) by:

$$\Psi = 9.8 \cdot 10^{-3} \frac{t_{AIR} R}{\rho_w g v_w} H_R, \quad (27)$$

with R being the gas constant ($R = 8.31 \text{ J mol}^{-1} \text{ K}^{-1}$), ρ_w and v_w the density and molar volume of liquid water (here considered as $\rho_w = 1000 \text{ kg m}^{-3}$ and $v_w = 1.8 \cdot 10^5 \text{ m}^3$

mol⁻¹), respectively, and g the gravitational acceleration ($g = 9.81 \text{ m s}^{-2}$). The atmospheric water potential was estimated by Eq. (27) using air relative humidity and temperature measured for each treatment every 30 min.

The mean soil water pressure head h_m (MPa) experienced by the root system of plants at each observation point (three depths at two observation points per irrigation treatment) was calculated using root distribution over depth as weighing factor. During the reproductive phase, the root system was supposed to be completely developed, and therefore we assumed a constant relative root length distribution over time. For our simulations, root length distribution for *Phaseolus vulgaris* L. as described by Guimarães *et al.* (1993) was used. His experimental work was done in climatic and soil conditions similar to our experiment and with the same crop cultivar, reporting almost 50% of the total root length in the upper 25% of the rooted profile (0-0.1 m), another 25% in the second 25% of the profile (0.1-0.2 m), while the bottom 50% of the profile accounted for approximately 25% of the total root length (0.20.4 m). In agreement with these measurements, we assumed h_m equal to:

$$h_{m,i} = 0.5h_{1,i} + 0.25h_{2,i} + 0.25h_{3,i}, \quad (28)$$

where: the subscript i indicates the observation point in the field (1 and 2 in the fully irrigated treatment, and 3 and 4 in the deficit irrigated treatment), and h_1 , h_2 and h_3 (MPa) are the soil water pressure heads measured at 0.05, 0.1 and 0.3 m depth, respectively.

Before applying a PCA, a detrending operation was applied to separate the simulated transpiration (T_{FI} and T_{DI}) data sets in temporal components. Detrending is the statistical or mathematical technique of removing trends from data sets (Moran *et al.*, 2009). Two main trends are usually observed in transpiration data: a daily trend correlated to the daily radiation cycle, and a long term trend correlated to soil water availability. Detrending was performed on the transpiration rate estimated every 30 min in daytime by the best performing Ag_s model (the model version showing the best statistical indices) to remove the long term trend. Initially, time-dependent polynomial equations were fitted to the transpiration rate data series between August 4 (day 0) and September 2 (day 29). For the fully irrigated treatment, a first-degree polynomial equation was used, and for the deficit irrigated treatment a third-degree polynomial equation showed the best performance. The long term transpiration rate T_{lt} was then calculated by the fitted polynomial equations for each irrigation treatment ($T_{lt, FI}$ and $T_{lt, DI}$):

$$T_{lt, FI} = a + b\tau, \quad (29)$$

and

$$T_{lt, DI} = c + d\tau + e\tau^2 + f\tau^3, \quad (30)$$

where: τ is the time (d) and a , b , c , d , e and f are fitting parameters.

For each treatment, the daily component of transpiration ($T_{daily, FI}$ and $T_{daily, DI}$) was calculated by:

$$T_{daily, FI} = (T_{FI} - T_{lt, FI}) + T_{min Tt, FI}, \quad (31)$$

and

$$T_{daily, DI} = (T_{DI} - T_{lt, DI}) + T_{min Tt, DI}. \quad (32)$$

The addition of the minimum values of T_{lt} to the difference forces transpiration rates to be positive, thus preserving their physiological meaning.

In addition to daily and long term trends in transpiration rates, a third trend that is frequently observed is a transpiration reduction during the warmest hours of the day, sometimes referred to as the afternoon depression of photosynthesis (Wang *et al.*, 2006). Even when soil water conditions are favourable, stomatal closure is observed due to high atmospheric demand for water vapour. Leuning (1995) and Tuzet *et al.* (2003) explained in details the afternoon depression of photosynthesis. According to those authors, stomatal conductance depends not only on net radiation, temperature and intercellular CO₂ concentration via photosynthesis, but also on leaf water potential, which is in turn a function of soil water potential and water flux rate between soil and plant. During the hours of very high atmospheric demand, the combined soil-plant resistance is too high to provide sufficient water, and a depression of photosynthesis can develop by stomatal closure. As a result, the canopy temperature at midday has a maximum value (Wang *et al.*, 2006). Using canopy temperature data measured by infrared thermometry as input in the Ag_s model, that uses the vapour pressure deficit between intercellular

air spaces of leaves and atmospheric air as the driving force for the diffusion of vapour, the afternoon depression of photosynthesis can be simulated (Jones, 2004) and included in estimated transpiration rates. Days that presented this feature in the T_{daily} component were selected and the simulated transpiration rates, identified as $T_{d, FI}$ and $T_{d, DI}$ for the fully and deficit irrigated treatments, respectively, were also analysed with the PCA methodology.

RESULTS AND DISCUSSION

To quantitatively evaluate the model performance with all parameterizations, statistical indices $RMSEP$, ME and E of estimated and observed canopy temperature t_c are presented in Table 3 for both irrigation treatments. Values of $RMSE$ and ME closer to zero indicate good model prediction compared to the observations, while the coefficient of efficiency (E) ranges from $-\infty$ to 1, with positive values indicating that model predictions are better than the simple data average. The most important error decrease in t_c estimation and an increase in model efficiency were observed when the g_m parameterization was modified. Changes in Γ , $A_{m, max}$ and R_d parameterizations did not effectively alter model performance when compared to the original model. The Amthor (1995) function for G is close to the original one and presented almost the same statistical performance, while the functions of Ballantyne *et al.* (2011) and Bonan *et al.* (2011) result in smaller values of Γ than the original function increased the error of prediction. This indicates that greater values of Γ with t_c better represent a water stress condition. The Flowers *et al.* (2007) parameterization for $A_{m, max}$ reduced prediction errors, indicating that the new function for $A_{m, max}$ with a low peak would better represent a condition of water stress. Inserting the parameterization

Table 3. Root-mean-square error of prediction $RMSEP$ (°C), mean error ME (°C) and coefficient of efficiency E of the canopy temperature estimated by the original Ag_s model by Jacobs (1994) and by the model with new parameterizations

Parameterization	$RMSEP$ (°C)		ME (°C)		E	
	FI	DI	FI	DI	FI	DI
Original Jacobs (1994) model	8.09	7.32	3.54	2.32	-1.24	0.1
Γ – Amthor (1995)	8.10	7.32	3.55	2.32	-1.24	0.1
Γ – Ballantyne <i>et al.</i> (2011)	8.21	7.41	3.62	2.38	-1.3	0.08
Γ – Bonan <i>et al.</i> (2011)	8.41	7.58	3.73	2.47	-1.42	0.04
g_m – Flowers <i>et al.</i> (2007)	5.91	6.43	-0.81	-1.28	-0.19	0.3
$A_{m, max}$ – Goudriaan and van Laar (1994)	7.85	7.19	3.25	2.12	-1.11	0.13
R_d - Goudriaan and van Laar (1994)	8.24	7.38	3.64	2.32	-1.32	0.09
G Planophile-Roujean (1996)	8.39	7.67	3.32	2.23	-1.41	0.02
G Erectophile-Roujean (1996)	8.08	7.26	3.64	2.35	-1.23	0.12

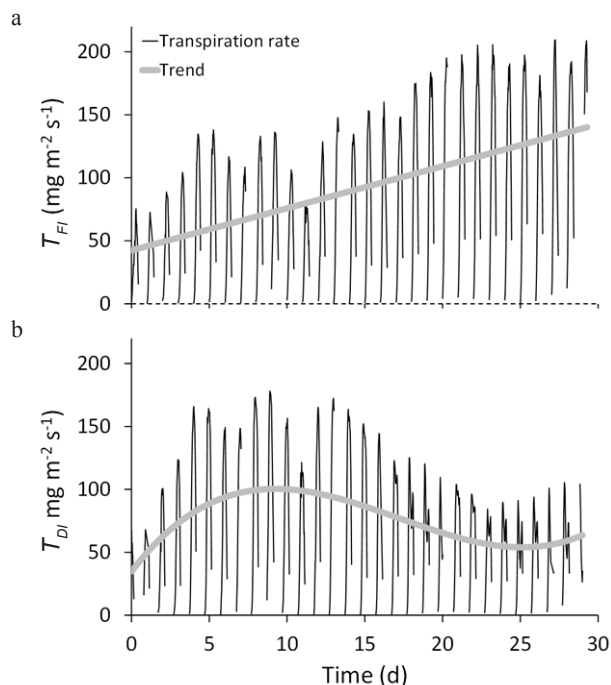


Fig. 2. Plant transpiration rate estimated by the A_g_s model for: a – the fully irrigated treatment (T_{FI} , $\text{mg m}^{-2} \text{s}^{-1}$), and b – the deficit irrigated treatment (T_{DI} , $\text{mg m}^{-2} \text{s}^{-1}$) which presented the best statistical indices ($g_{m,25} = 3.9 \text{ mm s}^{-1}$ for the FI and $g_{m,25} = 4.4 \text{ mm s}^{-1}$ for the DI) and the respective temporal trends (FI : $T_{FI} = 42.455 + 3.337 \tau$; $R^2 = 0.22$; $T_{DI} = 34.764 + 16.175 \tau - 1.196 \tau^2 + 0.0232 \tau^3$; $R^2 = 0.13$).

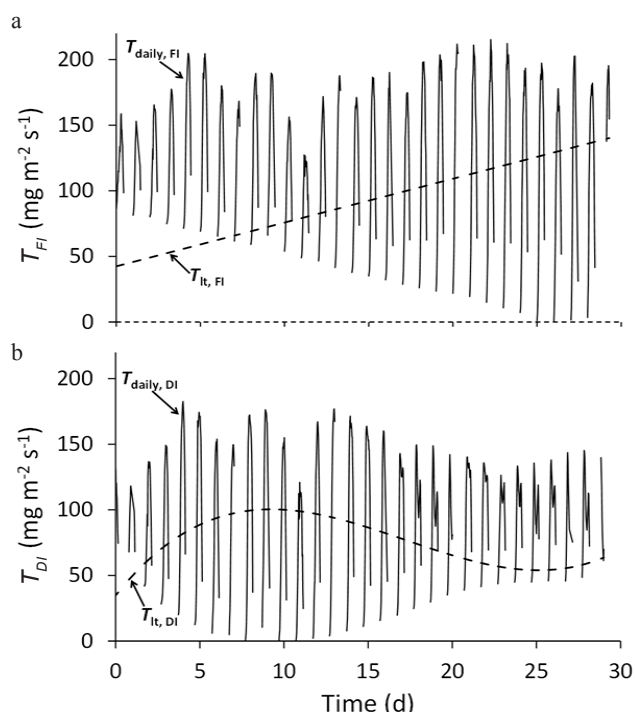


Fig. 3. Daily (T_{daily} , $\text{mg m}^{-2} \text{s}^{-1}$) and long term (T_{lt} , $\text{mg m}^{-2} \text{s}^{-1}$) components of plant transpiration rate for: a – the fully irrigated treatment ($T_{\text{daily,FI}}$ and $T_{\text{lt,FI}}$), and b – the deficit irrigated treatment ($T_{\text{daily,DI}}$ and $T_{\text{lt,DI}}$).

of R_d proposed by Goudriaan and van Laar (1994) in the A_g_s model increased prediction errors, indicating that the original parameterization for R_d could better represent the water stress condition. Modifying the leaf angle distribution G to erectophile as suggested by Roujean (1996) reduced t_c prediction errors and increased model efficiency, especially for the deficit irrigated treatment.

The model was shown to be most sensitive to g_m and a significant improvement in model performance was obtained by reducing the value of $g_{m,25}$ to 3.9 mm s^{-1} , as suggested by Flowers *et al.* (2007). We investigated the optimum value for $g_{m,25}$ and found that 3.9 mm s^{-1} is, in fact, the best value for the fully irrigated treatment (FI). For the deficit irrigated treatment (DI), $g_{m,25} = 4.4 \text{ mm s}^{-1}$ gave the best model performance ($RMSEP = 6.39^\circ\text{C}$, $ME = 0.6^\circ\text{C}$, $E = 0.32$). The diffusional limitations to photosynthesis imposed by reductions in mesophyll conductance in water limited environments is not reported to be important to all common bean cultivars (Lizana *et al.*, 2006). However, for the Carioca cultivar used in our study a reduced value of $g_{m,25}$ better represented the photosynthetic process when plants were water stressed.

The obtained transpiration rates with $g_{m,25} = 3.9 \text{ mm s}^{-1}$ for the FI treatment and with $g_{m,25} = 4.4 \text{ mm s}^{-1}$ for the DI treatment were plotted against time to correlate the estimated transpiration rate to different environmental-plant variables which control the transpiration process on different time scales (Fig. 2). The temporal trend during the entire growth cycle could be described by a first-degree polynomial equation for the FI treatment ($T_{FI} = 42.455 + 3.337 \tau$; $R^2 = 0.22$), and by a third-degree polynomial equation for the DI treatment ($T_{DI} = 34.764 + 16.175 \tau - 1.196 \tau^2 + 0.0232 \tau^3$; $R^2 = 0.13$). Well-watered bean plants of the FI treatment linearly increased transpiration rate, whereas plants under water stress (DI treatment) showed an increasing trend in transpiration rate for some days after irrigation had been suspended but it decreased with time as the soil dried out. This suggests that transpiration was limited by water shortage in the soil and in this case the atmospheric demand for water vapour did not affect the transpiration rate, as discussed by Tuzet *et al.* (2003) and Medina and Gilbert (2016). By the end of the period, an irrigation gift of $\sim 15 \text{ mm}$ was applied to guarantee plant survival and the trend of transpiration rate was to increase.

Using the trend equations, a detrending was performed on half-hourly transpiration rate data series in daytime to separate it in the long term (T_{lt}) and daily (T_{daily}) temporal components for both irrigation treatments (Fig. 3). The T_{lt} component was obtained by the fitted polynomial equations and T_{daily} was given by taking the difference between the T estimates of the A_g_s model and T_{lt} and adding the minimum values of T_{lt} : $98.5 \text{ mg m}^{-2} \text{s}^{-1}$ for the deficit irrigated treatment, and $126.68 \text{ mg m}^{-2} \text{s}^{-1}$ for the fully irrigated treatment. The afternoon depression (T_d), the third temporal component of transpiration data sets, could also be identified in

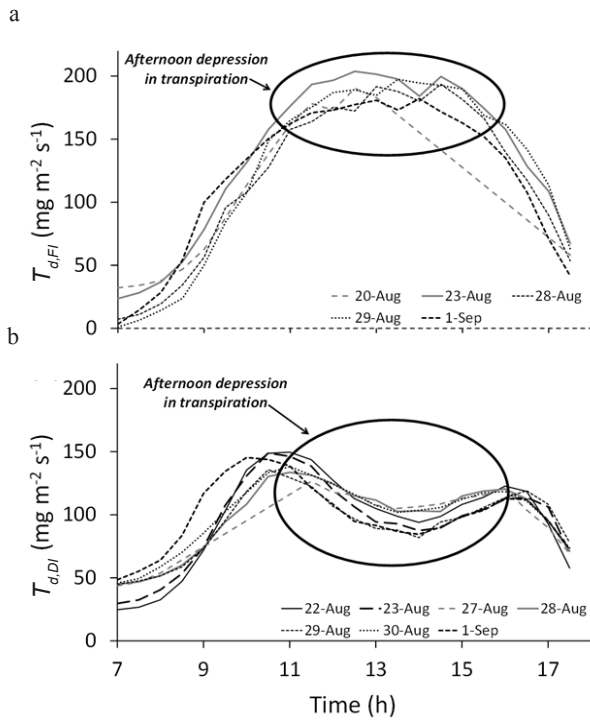


Fig. 4. Afternoon depression (T_d , $\text{mg m}^{-2} \text{s}^{-1}$) component of plant transpiration rate for: a – the fully irrigated treatment (T_{dFI}), and b – the deficit irrigated treatment (T_{dDI}).

T_{daily} data sets (Fig. 4). It shows a noise in transpiration rates between 11 h and 16 h on 5 days in the FI treatment and a deep depression in transpiration around this period on 7 days in the DI treatment. As can be seen in Fig. 4, the depression in transpiration of well-watered plants is much less pronounced than the one observed in water stressed plants.

After separating the transpiration data series in three main temporal components (T_{daily} , T_{lt} and T_d), each one was correlated to environmental and plant forcing variables. Temporal components were analysed using the principal component analysis (PCA) methodology. Prior to PCA, the atmospheric water potential ψ was calculated using Eq. (27) for both irrigation treatments and is presented in Fig. 5 together with the mean soil water pressure head h_m . While the soil became much drier in the deficit irrigated treatment than in the fully irrigated treatment (minimum $h_{m,3}$ and $h_{m,4} \approx -0.49$ MPa, and minimum $h_{m,1}$ and $h_{m,2} \approx -0.03$ MPa), ψ decreased almost equally in both of them as there was no rainfall for 53 days and the atmospheric potential decreased. Therefore, although plants in the FI treatment grew in a well-watered soil, they were subjected to a very dry atmospheric environment. High vapour pressure deficit in the atmospheric air is reported to decrease stomatal conductance and transpiration rate (Bunce, 1997; Shekoofa *et al.*, 2016) even in a well-watered soil.

The PCA was performed by taking the data series of T_{daily} , T_{lt} and T_d together with selected environmental-plant variables (identified in the following by EP): specific

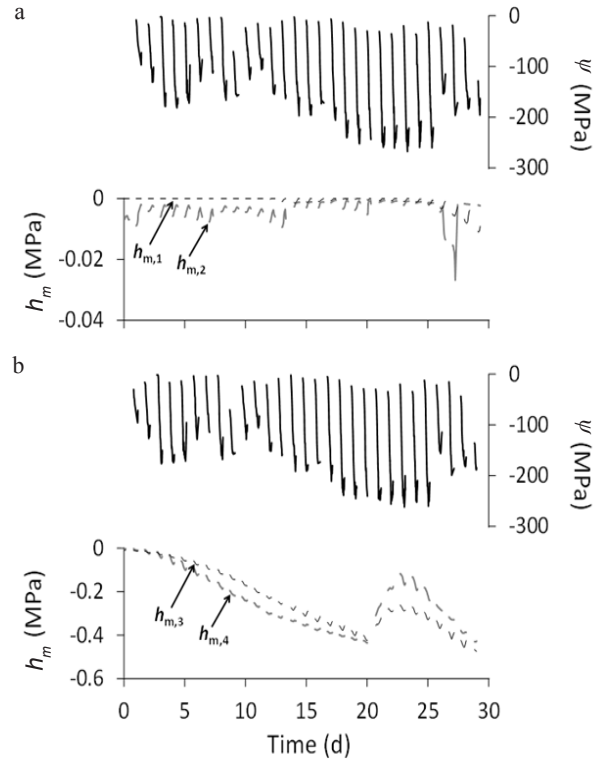


Fig. 5. Mean soil water pressure head h_m (MPa – Eq. (28)) at two observation points and the atmospheric water potential ψ (MPa) observed in: a – the fully irrigated treatment (FI), and b – the deficit irrigated treatment (DI).

humidity deficit between leaf and air (D_s), photosynthetically active radiation (I_{PAR}), relative humidity (H_R), canopy and air temperature (t_c and t_{air}), wind speed (u), mean soil water pressure head at two observation points each irrigation treatment (h_m), atmospheric water potential ψ , and leaf area index (L_{LAI}). To characterise the environment which was controlling plant transpiration, PCA was also performed by taking only the selected environmental and plant variables (EP for T_{daily} , T_{lt} and T_d). All accumulated variances of the principal components are presented in Table 4. For both FI and DI treatments the first three principal components accounted for more than 80% of the variance of the data sets of estimated transpiration rates combined with the data sets of the ten environmental-plant variables (totalling eleven variables). The same was observed when only environmental-plant data sets (ten variables) were analysed. In all cases, the first principal component accounted for more than 50% of the variance in data sets. In this case, those variables cumulatively accounting for approximately 20% of total variance can be removed from the data set with a negligible loss of information (Mundlak, 1981). Data of the environmental-plant data sets for T_{daily} and T_{lt} components are different from those of the T_d component as T_d represents only some days of the entire T_{daily} data sets. For this reason, they are presented separately in Tables 4-5 (named as EP for $T_{daily, FI}$ and $T_{lt, FI}$, EP for $T_{daily, DI}$ and $T_{lt, DI}$, EP for $T_{d, FI}$ and EP for $T_{d, DI}$).

Table 4. Accumulated variance explained by the principal components related to the daily, long term, and depression transpiration time components and related to the environmental-plant variables (EP) for the fully irrigated treatment (T_{daily}^{FI} , T_{lt}^{FI} , T_{d}^{FI} , EP for T_{daily}^{FI} and T_{lt}^{FI} , and EP for T_{d}^{FI}), and deficit irrigated treatment (T_{daily}^{DI} , T_{lt}^{DI} , T_{d}^{DI} , EP for T_{daily}^{DI} and T_{lt}^{DI} , and EP for T_{d}^{DI}). Variance could be explained by the principal component 1, 2 and 3 in both irrigation treatments

Irrigation treatment	Components for PCA	Accumulated variance (%)		
		Principal component		
		1	1 and 2	1, 2 and 3
FI	T_{daily}^{FI}	54.9	70.2	82.7
	T_{lt}^{FI}	50.1	70.3	82.6
	EP for T_{daily}^{FI} and T_{lt}^{FI}	52.8	68.5	82.0
	T_{d}^{FI}	57.5	75.9	88.6
	EP for T_{d}^{FI}	54.6	74.8	87.9
DI	T_{daily}^{DI}	54.0	74.1	84.6
	T_{lt}^{DI}	50.8	71	84.1
	EP for T_{daily}^{DI} and T_{lt}^{DI}	54.7	74.4	85.8
	T_{d}^{DI}	52.2	76.0	86.7
	EP for T_{d}^{DI}	51.8	78.0	89.0

From the foregoing we find that the transpiration rate data sets and environmental and plant data sets can be represented by three components. The next step is to identify which three variables can best represent these data sets. In Table 5 we present the percentage of variance of each input variable accounted for (final communality) for *FI* and *DI* treatments, respectively. Unlike the cumulative proportion, these statistics are related to one input variable at a time and it is possible to detect which input variables are better or worse represented by the reduction of the number of parameters.

For the fully irrigated treatment (Table 5), the specific humidity deficit between stomata and air (D_s), the relative air humidity (H_R) and the canopy temperature (t_c) are the variables that capture most of the variance on daily and long term temporal components of transpiration and environmental-plant data sets. This is in agreement with Fletcher *et al.* (2007), who reported close correlation between transpiration rate of well-watered soybean plants and vapour pressure deficit, a variable representing the atmospheric demand to water vapour like D_s and H_R . For the component T_{d}^{FI} the variance of H_R slightly decreased and the third most representative variable was the estimated transpiration rate itself. This is explained by the fact that the afternoon transpiration depression occurs when plants close their stomata (increasing t_c) in response to the high midday atmospheric demand to water vapour (high D_s and low H_R) (Leuning, 1995; Hérould *et al.*, 2013). As soil water was not limiting

plant growth in this treatment, simulated plant transpiration mainly responded to the atmospheric controlling variables, irrespective of the temporal component.

In the *DI* treatment (Table 5), the variables representing the variance in the data sets changed in relation to the *FI* treatment and also between temporal components. Relative air humidity (H_R), canopy temperature (t_c) and mean soil water pressure head at observation point 3 ($h_{m,3}$) are the most representative variables for the T_{daily}^{DI} component, while t_c and the mean soil water pressure head at points 3 and 4 ($h_{m,3}$ and $h_{m,4}$) describe an important part of the variance in the T_{lt}^{DI} component. Although representing a major part of the variance in both data sets (T_{daily}^{DI} and T_{lt}^{DI}), as water shortage in the soil was limiting transpiration, mean soil water pressure head accounted for a large part of the variance in the T_{d}^{DI} component. The environmental-plant data set (EP) also showed a high influence of H_R , t_c and $h_{m,3}$ on its main variance. These variables are more related to a plant water stress condition due to a dry atmospheric air, a high sensitive heat load on plants, and a soil water shortage, respectively, which are conditions likely to occur in the deficit irrigated treatment. The T_{d}^{DI} data set and the environmental-plant data set could be mainly represented by H_R , t_c and $h_{m,4}$. The afternoon depression in transpiration of the *DI* treatment occurred when plants closed their stomata (increasing t_c) in a combined response to the high atmospheric demand to water vapour (low H_R) and to the soil water content shortage (very low $h_{m,4}$) (Martínez-Vilalta *et al.*, 2014).

Table 5. Final communality (%) for the variables most related to the principal components for transpiration simulation of the fully and deficit irrigated treatments (*EP*: environmental-plant variables; D_s : specific humidity deficit between leaf and air; I_{PAR} : photosynthetically active radiation; H_R : relative humidity; t_c : canopy temperature; t_{air} : air temperature; u : wind speed; h_m : mean soil water pressure head; ψ : atmospheric water potential; L_{LAI} : leaf area index)

Parameter	Final communality (%)				
	$T_{daily, FI}$	$T_{lt, FI}$	<i>EP</i> for $T_{daily, FI}$ and $T_{lt, FI}$	$T_{d, FI}$	<i>EP</i> for $T_{d, FI}$
T_{daily} , or T_{lt} , or T_d ($\text{mg m}^{-2} \text{s}^{-1}$)	91.6	92.8	-	96.4	-
D_s (mg kg^{-1})	97.7	96.9	96.9	98.8	98.4
I_{PAR} (W m^{-2})	62.7	58.3	57.7	74.2	70.2
H_R (%)	94.3	95.4	95.4	95.2	96.1
t_c ($^{\circ}\text{C}$)	94.6	94.5	94.5	97.2	97.2
t_{air} ($^{\circ}\text{C}$)	88.7	89.2	89.5	94.8	95.1
u (mm s^{-1})	58.8	59.7	66.0	77.5	80.0
$h_{m,1}$ (MPa)	84.6	74.9	85.7	91.8	92.0
$h_{m,2}$ (MPa)	63.1	63.3	63.2	66.4	67.3
ψ (MPa)	89.3	90.1	89.9	89.2	89.6
L_{LAI} ($\text{m}^2 \text{m}^{-2}$)	83.9	93.3	81.5	92.7	92.7
	$T_{daily, DI}$	$T_{lt, DI}$	<i>EP</i> for $T_{daily, DI}$ and $T_{lt, DI}$	$T_{d, DI}$	<i>EP</i> for $T_{d, DI}$
T_{daily} , or T_{lt} , or T_d ($\text{mg m}^{-2} \text{s}^{-1}$)	78.1	84.7	-	70.1	-
D_s (mg kg^{-1})	92.7	94.4	94.6	92.6	95.2
I_{PAR} (W m^{-2})	60.0	47.9	55.3	67.5	64.0
H_R (%)	95.6	95.3	94.9	98.2	97.3
t_c ($^{\circ}\text{C}$)	96.6	97.2	97.2	97.5	98.3
t_{air} ($^{\circ}\text{C}$)	91.5	90.7	92.4	94.4	94.3
u (mm s^{-1})	67.0	40.4	72.4	73.3	79.6
$h_{m,3}$ (MPa)	99.0	99.0	99.4	82.8	83.3
$h_{m,4}$ (MPa)	90.7	98.0	90.9	99.6	99.6
ψ (MPa)	88.4	86.3	89.4	95.0	94.8
L_{LAI} ($\text{m}^2 \text{m}^{-2}$)	70.9	91.5	71.9	82.5	84.1

CONCLUSIONS

1. Using the parameterizations recovered in the Ag_s model and simulating plant transpiration rate resulted in a moderate reduction of prediction errors and in an increase in model performance. The model was shown to be the most sensitive to changes in the mesophyll conductance parameterization, and special attention should be given to the parameterization for the leaf angle distribution when

transpiration of plants under water stress is estimated. The use of an erectophile function for leaf angle distribution, better representing leaf distribution under water stress, also improved model performance.

2. Transpiration data sets simulated with a new parameterization for mesophyll conductance could be separated in the temporal components which were well related to the main forcing variables controlling plant transpiration for each environmental and soil water condition. The daily, the

afternoon depression and the long term transpiration temporal components of the fully irrigated treatment were more related to the atmospheric forcing variables which were causing a high transpiration rate due to the dry atmospheric air (stomata-air specific humidity deficit, relative air humidity and canopy temperature). The daily and afternoon depression transpiration temporal components of the deficit irrigated treatment were related to both the atmospheric and soil dryness, whereas the long term transpiration component of this treatment mostly responded to soil water shortage. These qualitative relations were determined by applying the principal component analysis methodology.

3. The modified versions of the A_g_s model by Jacobs (1994) were able to simulate plant transpiration and its three temporal components in very dry conditions. Without using soil data as a model input, consistent values of transpiration in a temporal component closely related to soil water availability (long term) were simulated by a parameterization representative of the water deficit condition. These results show the ability of the A_g_s model to simulate plant transpiration under dry conditions without using specific soil input data, thus avoiding the need to obtain soil data and to consider the naturally occurring soil heterogeneity.

Conflict of interest: The Authors do not declare conflict of interest.

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