

Intersecting-surfaces approach to soil structure

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A b s t r a c t. The large number of intersecting surfaces in a volume divides the latter into small sub-volumes. Each of the sub-volumes is outlined or nearly outlined by parts of the intersecting surfaces. A number of simple geometrical conditions to be imposed on the intersecting surfaces determine a certain size distribution of the outlined sub-volumes. The approach based on these prerequisites was previously successfully applied to the multiple cracking and fragmentation of rocks, naturally and by blasting, as well as to soil fragmentation by shrinkage cracks. In both cases the intersecting surfaces are cracks, and sub-volumes are rock or soil fragments. The first application of the approach to the pore structure of soils was related to modelling the clay matrix pores and their size distribution. In this case the plate-like clay particles play the part of the intersecting surfaces, and the pore volumes play the part of the sub-volumes. The approach leads to the scale-invariant fragment- and pore-size distributions and permits one to take into account the superposition of the intersecting surfaces of a different physical nature and scale, the connectedness and tortuosity of the intersecting surfaces, the shape characteristics of sub-volumes, and the swelling-shrinkage of a system of clay particles. In general, soils contain, besides clay matrix pores, such sub-volumes as silt-sand grains and different types of clay aggregates and pores. The boundaries of the silt-sand grains and aggregates, along with cracks and clay particles, play the part of the intersecting surfaces. The objective of this work is to generalize the approach to model the different size distributions in sand and aggregated soil (in the last case we only consider the interaggregate, *ie* structural porosity). This development follows a brief summary of the approach including its previous applications as well as a consideration of the differences between such intersecting surfaces as cracks, plate-like clay particles, boundaries of sand grains and aggregates. A comparison between the model predicted size distributions and relevant available data shows satisfactory agreement.

K e y w o r d s: surface- and volume-like objects, pores, grains, aggregates

INTRODUCTION

Soil structure plays a key part in modelling different properties of the soil including those which are mechanical, hydraulic, and shrink-swell. Quantitatively, the soil structure is characterized by size and shape distributions of different pore types, size and shape distributions of different solids forming the soil, as well as connectedness and tortuosity of pore walls and channels. Geometrically, soils consist of volume-like objects (V-objects – blocks, fragments, aggregates, silt and sand grains, and 3D pores) and surface-like objects (S-objects – boundaries of V-objects as well as cracks, slits and clay particles). In the following this aspect will be discussed in more detail.

There are a number of approaches to modelling pore-size distributions (as a part of functions describing soil structure). The statistical approach has been used to represent the distribution of the pore and aggregate sizes. For instance, Gardner (1956) and, more recently, Kosugi (1994, 1996) considered a lognormal distribution as a convenient mathematical representation to be justified empirically. In the frame of this approach there are also other probability laws (Brutsaert, 1966; Rice, 1995). Arya and Paris (1981) constructed the pore size distribution based on already known grain size distribution. The approach includes a number of assumptions relative to pore characteristics such as the length of a pore and pore opening. Rieu and Sposito (1991a, b) proposed a fractal model of soil porosity of aggregated soils including pores and aggregates of only one type. Nimmo (1997) introduced a division of pore space into 'texture' and 'structure' ranges. For testing the model predicted pore size distributions one usually uses data of the soil water retention curve (Perrier *et al.*, 1996) or the data of

porosimetry (Newman and Thomasson, 1979; Fies and Bruand, 1990). Interpretation of such data includes an assumption about effectively cylindrical pores and their filling in up to a certain radius.

It is worth noting that the above approaches only deal with pore and aggregate size distributions, *ie* they come to a non-total set of expressions from the viewpoint of the quantitative aspect of soil structure *eg* pore shape distributions are not considered. This remark, the above mentioned understanding of soil structure, and analysis of indicated approaches enable one to formulate the following conditions, the observance of which is desirable at the derivation and validation of expressions quantitatively describing the soil structure.

1. *S- and V-objects as a basis.* It is preferable to construct different size distributions of a soil structure based on physical considerations, but not simply as convenient empirically justified mathematical representations. Below we refer to such a distribution as the physical one. One of the major (and obvious) physical factors which should be taken into account is the co-existence in a soil of the networks of S-objects and corresponding systems of V-objects (of different scale and nature).

2. *V-objects of a different type on equal rights.* Pore size distributions should be constructed not based on already known grain (or aggregate) size distribution but jointly with the latter. That is, a model should describe the size distributions of V-objects of different types on equal rights and for the type number equal to or more than two, *eg* sand grains, clay aggregates, and pores.

3. *The modes of V-objects of a given type.* V-objects of a given type, *eg* pores, can include a number of modes with different maximum dimensions. One can consider the 'texture' and 'structure' ranges of a pore space (Nimmo, 1997) as a particular case of such modes

4. *Distribution validation of solids is simultaneously that of corresponding pores.* For validation of a model predicted pore size distribution along with water retention and porosimetry data one can use the validation of the corresponding model predicted aggregate and grain size distributions based on the data of the distributions. Such validation is also direct because theoretically these distributions are obtained jointly (see condition 2).

In light of these conditions one can assume that the geometrical aspect of a soil structure connected with S- and V-objects (see above) should play a key part. An approach for modelling different fragment and pore size distributions in soils and rocks was initially applied to cracks and clay particles as S-objects and to blocks and clay matrix pores as V-objects (Chertkov, 1986; 1991; 1995; 2000; 2003; Chertkov and Ravina, 1998; 1999). For the reader's convenience, a brief summary of the intersecting-surfaces approach (ISA) is given below (based on indicated works). The gist of the approach is as follows. The large number of

intersecting surfaces (S-objects) in a volume divides the latter into small sub-volumes (V-objects). Every one of the sub-volumes is outlined or nearly outlined by parts of the intersecting surfaces. A certain size distribution of the outlined V-objects originates from a number of simple geometrical conditions imposed on the S-objects.

The objective of this work is to generalize the intersecting-surfaces approach and, using the generalized ISA, to model the different size distributions in a sand and aggregated soil. The generalization should account for the simultaneous presence in a soil volume of a number of types of S- and V-objects and a number of modes of S- and V-objects of a similar type. Except for this novelty, the applications of the generalized ISA for sand and aggregated soil are new since they are based on considering the S-objects of another nature than cracks and clay particles (see 'Peculiarities of different S- and V-objects in soils'). Note that the aggregated soil is considered here at a given water content, *ie* shrinkage-swelling is beyond the scope of this work.

The development of the applications for a sand and aggregated soil follows a summary of major concepts and relations of ISA as well as brief consideration of differences between such S-objects as cracks, plate-like clay particles, boundaries of sand grains and aggregates. This consideration aims to preliminarily show the place of ISA applications to be developed in this work relative to existing ISA applications. After the development of the new applications, the model predicted size distributions of pores, sand grains, and aggregates are compared to relevant available data.

MAJOR CONCEPTS AND RELATIONS OF ISA
(Brief summary from Chertkov (1986, 1991, 1995, 2000, 2003),
Chertkov and Ravina (1998, 1999))

The concept of S- and V-objects

To define the three sizes of a V-object we consider the two sizes of its largest face and the size normal to the latter, and we designate them in order of decreasing values as $x \geq y \geq z$. By definition the size distribution of V-objects is the distribution with respect to their maximum size, x .

S-objects have limited sizes and random orientations in space. The ratio of connected S-objects to the total number of S-objects defines the so-called 'connectedness' of their system, $0 < C \leq 1$. S-objects do not stick together at their surfaces. Connections of S-objects of close sizes, $\sim x$, lead to the development of an S-object of a larger scale $\sim (K^*+1)x$, according to the concentration criterion at $K^* \cong 5$ (Zhurkov *et al.*, 1981). The condition of the effective independency of S-objects: the spacing x between neighboring intersections of the S-objects with a straight line is distributed as $\exp(-x/d)$ where: d is an average value (Hudson and Priest, 1979 – data for rocks; Scott *et al.*, 1986 – data for soils).

The condition of V-object formation at S-object connection: at least part of the S-objects outline V-objects totally or nearly totally. The V-objects have sizes in a range, $0 < x \leq x_m$ and constitute a volume fraction $f_m \leq 1$ of the soil (rock) volume.

The size distribution of V-objects

If the S-objects fulfill the above geometrical conditions, the physical size distribution of V-objects (see Introduction, condition 1) is assumed to originate from the division of a volume by a large number of intersecting S-objects. Then, the size distribution of V-objects, $F(x/x_m)$ (a volume fraction of V-objects of size $< x$ of their total volume) is given as:

$$F(x/x_m) = f(x/x_m) / f_m, \quad 0 < x \leq x_m \quad (1)$$

$$f(x/x_m) = 1 - \exp(-I(x/x_m)), \quad 0 < x \leq x_m \quad (2)$$

$$I(x/x_m) = C \ln(K^* + 1) (4x/x_m)^4 \exp(-4x/x_m), \quad (3)$$

$$x_m = 4d, \quad (4)$$

and

$$f_m \equiv f(1) \equiv 1 - \exp(-8.4C) \quad (5)$$

(the coefficient 8.4 originates from Eqs (2) and (3) at $x = x_m$ and $K^*=5$), $I(x/x_m)$ is the average S-object number of x size in $\sim x^3$ volume, $f(x/x_m)$ is the volume fraction of V-objects of size $< x$ of the soil (rock) volume, and f_m is the volume fraction of V-objects of the soil (rock) volume. If $C=1$, the entire soil (rock) volume is divided by S-objects into V-objects (the case of $f_m = 1$). If $C < 1$, the soil volume is not completely divided by S-objects ($f_m < 1$). There are also similar expressions for the 2D case.

The maximum size of V-objects, x_m and connectedness, C of S-objects, *eg* of a crack network, are two independent parameters of the distribution. These parameters can depend on variables specific for a particular system of S-objects and on spatial coordinates.

The features of the intersecting-surfaces approach

The characteristic dimension x_m simultaneously plays the part of the scaling factor as the denominator of the x/x_m ratio (Eqs (1)-(3)). That is, for different scales the form of the distribution (Eqs (1)-(5)) does not change. In this meaning the distribution is scale-invariant.

Connected S-objects outline V-objects. Hence, the F function and the shape of V-objects being known, one can estimate the *specific surface area*, L_3 , of connected S-objects (per unit volume of the soil). The specific length,

L_2 , of connected traces of the S-objects at a cross-section can be similarly estimated for the 2D case.

3D tortuosity, T_3 of intersecting (or connected) S-objects along a given direction is defined as a ratio:

$$T_3 = L_3 / L_2. \quad (6)$$

The tortuosity is expressed in terms of the $f(x/x_m)$ function and its 2D version. For the total S-object system (when $x = x_m$) the connectedness and tortuosity of the system are interconnected.

In connection with the soil heterogeneity and variability of a different type, x_m and C parameters become functions of spatial coordinates and time (*locality*).

The approach enables one to consider an actual soil as a *superposition* of S-objects of a different physical nature, *eg* clay particles, and cracks, and a different scale, *eg* cracks of essentially different sizes.

The probability $P(x, y, z) dx dy dz$ that a V-object will have three sizes in intervals dx, dy, dz , eventually flows out of the same regularity ($f(x/x_m)$):

$$P(x, y, z) = \frac{1}{f_m} \frac{df(x/x_m)}{dx} \frac{1}{f(x/x_m)} \frac{df(y/x_m)}{dy} \frac{1}{f(y/x_m)} \frac{df(z/x_m)}{dz}. \quad (7)$$

In the simplest case the restrictions imposed on the three sizes of a V-object have the form:

$$z \leq y \leq x \leq x_m. \quad (8)$$

The P function being known, one can estimate any *shape characteristic* of V-objects.

Thus, the approach leads to the scale-invariant fragment and pore size distributions and permits one to take into account the superposition of the S-objects of different physical nature and scale, the connectedness and tortuosity of the S-objects, the shape characteristics of V-objects, and the swelling-shrinkage of a system of clay particles.

Additional remarks

The concept of S- and V-objects is only connected with geometrical conditions. Therefore, systems of S-objects and corresponding V-objects meeting these conditions can have different dynamic properties, *eg* natural cracks and blocks in rock (Fig. 1a), shrinkage cracks and soil fragments (Fig. 1b), and plate-like clay particles and pores of clay matrix (Fig. 1c). The dynamic properties of a particular system of S- and V-objects can influence values of x_m and C parameters and their dependences on variables that are specific for the application, but not the form of the dimension distribution of the V-objects given by Eqs (1)-(5).

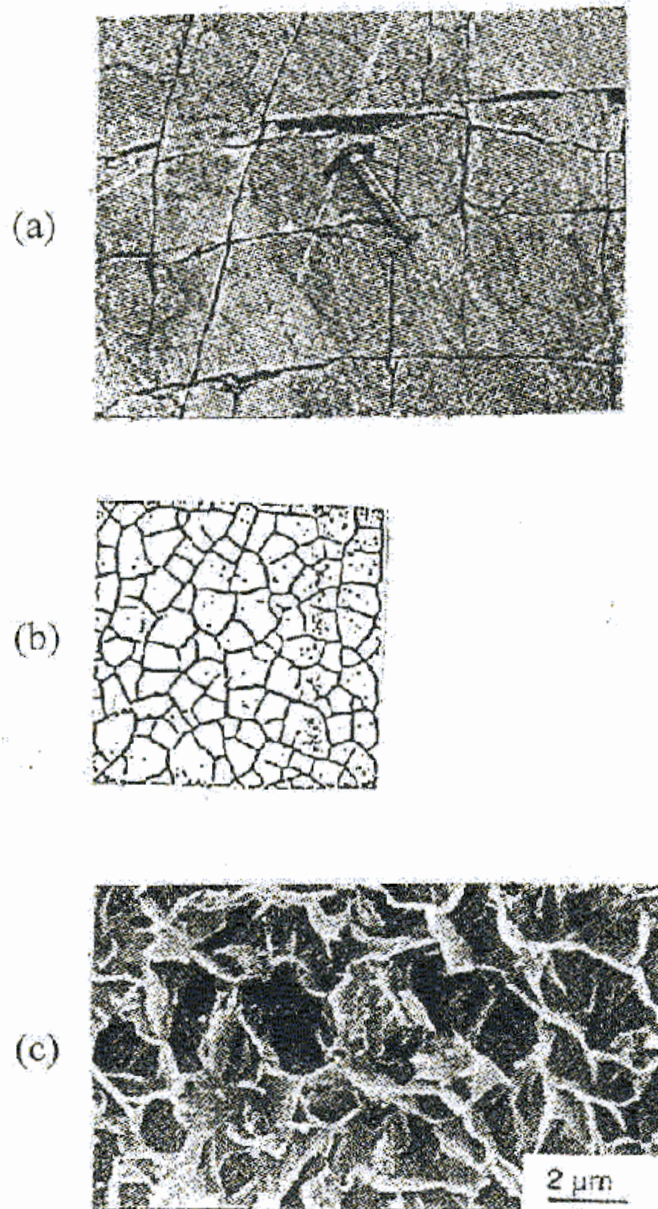


Fig. 1. Examples of real S-objects and corresponding V-objects of essentially different physical nature, scale, and dynamic properties, but meeting the similar ISA geometrical conditions (see the Section: The concept of S- and V-objects). a) Metamorphic rock with crack network (Verhoogen *et al.*, 1970 (Figs 3-18)). b) Sample of Fluvisol after shrinkage (Guidi *et al.*, 1978); the mean distance between cracks is about 6.5 mm. c) Microstructure of Greek Na-smectite (Wilding and Tessier's, 1988 (Fig. 16A)).

Part of the actual S-objects (cracks, slits, and clay particles) has small, but finite (non-zero) thickness (or aperture) and volume. Corresponding modifications can be introduced to the dimension distribution of V-objects given by Eqs (1)-(5) if the small thickness (aperture) is found by an independent way.

Existing applications

The intersecting-surfaces approach (ISA) based on these general geometrical prerequisites was applied to

multiple cracking and fragmentation of rocks (Chertkov, 1986, 1991), as well as to soil fragmentation by shrinkage cracks (Chertkov, 1995; Chertkov and Ravina, 1998). In both cases the intersecting surfaces (S-objects) are cracks, and V-objects are rock or soil fragments. As an example, let us consider the soil fragmentation by shrinkage cracks. The data relate to clod size distribution in artificially dried thin soil samples of Italian Fluvisol clay soil (Fig. 1b) from Guidi *et al.* (1978). The soil samples were formed by stacking aggregates of the 1-2 mm fraction of the original soil. The

method of clod size distribution measurement (Guidi *et al.*, 1978) was based on electro-optical determinations. The experimental data (histogram in Fig. 2) are shown together with the theoretical size distribution (solid curve in Fig. 2) corresponding to connectedness $C=1$ and the mean clod size $d=6.62$ mm (Eqs (1)-(5)). Agreement between the theoretical curve and the data is quite satisfactory.

The first application of the approach to soil structure related to modelling the clay matrix pores and their size distribution (Chertkov, 2000; 2003). In this case plate-like clay particles play the part of the intersecting surfaces (S-objects), and the pore volumes play the part of V-objects. The data relate to the pore size distribution (squares and triangles in Fig. 3) in Macon kaolinite, Georgia (Diamond,

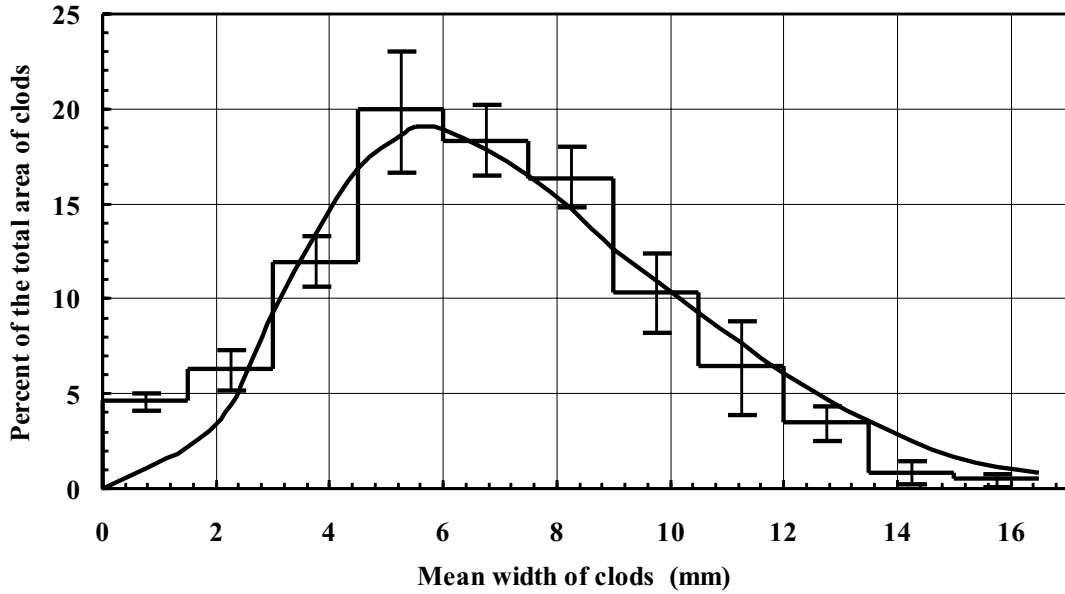


Fig. 2. Histogram of area distribution of 2D clods in the Fluvisol samples (Guidi *et al.*, 1978) and theoretical solid curve (Chertkov, 1995).

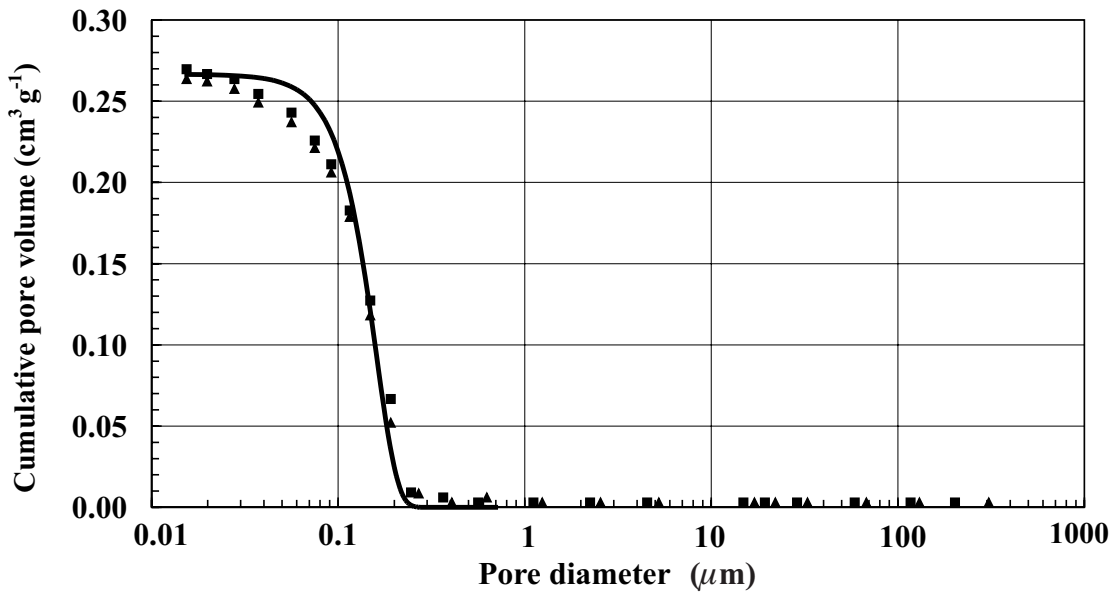


Fig. 3. The predicted pore-size distribution (Chertkov, 2000) and two data sets from Diamond (1970) for Macon, Georgia kaolinite.

1970). Pore size distribution (after oven drying) was studied by the mercury intrusion method. Model predicted distribution (solid curve in Fig. 3) corresponding to the connectedness of clay particles, $C=1$ and the maximum diameter of pore tubes, $\sim 0.72 \mu\text{m}$ (Eqs (1)-(5)) was found by the least-square criterion. Agreement between the data and model is also satisfactory.

PECULIARITIES OF DIFFERENT S- AND V-OBJECTS IN SOILS

The above existing applications relate to systems of S-objects of two types: cracks (of different origins) in soil (rock) and clay particles. In general, soils, besides fragments outlined by cracks and clay matrix pores outlined by plate-like clay particles, contain such V-objects as silt-sand grains, clay aggregates, and pores of different types. The boundaries of the silt-sand grains, aggregates, and pores of different types along with cracks and clay particles, play the part of S-objects. Different S-objects (and corresponding V-objects) in soils have a number of peculiarities that are important from the viewpoint of ISA applications.

1. Cracks, unlike other S-objects, can appear and enlarge. Other S-objects exist retaining their size. Accordingly, soil (rock) fragments can divide into smaller fragments (see Fig. 1a and b). (Size changes by swelling-shrinkage are not considered here).

2. Cracks and clay particles, unlike other S-objects, outline V-objects of a single type. Cracks outline soil (rock) fragments (Fig. 1a and b). Clay particles outline clay matrix pores (Fig. 1c). In other words, every crack (clay particle) is between two fragments (clay matrix pores). For this reason, soil (rock) fragments and clay matrix pores (unlike other V-objects) can form the total V-object system, *ie* fragments can occupy the total fragmented volume, and clay matrix pores can occupy the total clay matrix volume (see section 'Additional remarks' for a reservation about the non-zero thickness of clay particles).

3. Boundaries of soil aggregates outline two types of V-objects: aggregates and interaggregate (or structural) pores. The total system of the V-objects (the total volume) includes both aggregates and structural pores (unlike the case of soil fragments and clay matrix pores). Boundaries of structural pores separate the pores of aggregates. The surface area of the boundaries of structural pores is always a part of the surface area of aggregate boundaries. Therefore, the formal transition in the model to the limit $P \rightarrow 0$, where: P is structural porosity is possible unlike to the limit $P \rightarrow 1$.

4. Boundaries of sand grains also outline two types of V-objects: sand grains and pores of a sand skeleton. All that was said in point 3 about the soil aggregate system is the same for the sand grain system after changing "aggregates" by 'sand grains' and 'structural pores' by 'pores of sand skeleton'.

5. From the viewpoint of texture, clay matrix and sand skeleton are different structures. However, unlike soils of more complex texture and structure, their pore systems are the simplest because they include only one type of pores, clay matrix pores and sand skeleton pores, respectively. This feature also relates to the aggregated soil that contains only structural pores (here aggregates are considered to be 'structureless', *ie* sand grains, clay particles and matrix pores inside aggregates are ignored). The essential difference from the ISA viewpoint between a clay matrix on the one hand and sand and aggregated soil on the other, is as follows. Solids of a clay matrix are plate-like clay particles, *ie* S-objects. Solids of sand and aggregated soil are sand grains and aggregates, respectively, *ie* V-objects.

THE ELEMENTARY SIZE DISTRIBUTION OF V-OBJECTS

The V-object size distribution to be expressed by Eqs (1)-(3) and (5) can be written as:

$$F(x/x_m, P) = \frac{1 - (1 - P)^{I_o(x/x_m)/8.4}}{P}, \quad 0 < x \leq x_m \quad (9)$$

where: $I_o(\xi)$ is the function from Eq. (3) at $C=1$:

$$I_o(\xi) = \ln(K^* + 1)(4\xi)^4 \exp(-4\xi), \quad (K^* = 5) \quad (10)$$

and $P \equiv f_m$ is a volume fraction of V-objects. P is coupled with connectedness C from Eq. (5):

$$P \equiv f_m \cong 1 - \exp(-8.4C). \quad (11)$$

In the presentation of Eq. (9) the V-object size distribution is characterized by x_m and P parameters instead of x_m and C . We will refer to the distribution of Eq. (9) as an elementary ISA distribution of the V-objects of a given type. Figure 4 shows the numerical presentation of the elementary ISA distribution at different P values.

It is worth reiterating (see Section: Additional remarks) that the elementary ISA distribution of V-objects, *ie* objects that fulfill the abovementioned geometrical conditions, is a universal function of V-object size, x/x_m (Fig. 4). The view of the function (Eqs (9) and (10)) does not depend on V-object type, *eg* soil (rock) fragments or clay matrix pores (Fig. 1), and their methods of preparation. Materials and methods can only influence values of the P (or C) and x_m parameters.

GENERALIZATION: THE CASE OF SEVERAL TYPES OF S-OBJECTS AND CORRESPONDING V-OBJECTS

Initially (see Brief summary of ISA) the elementary distribution (Eqs (9) and (10)) was derived for the case where there is only one system of S-objects (cracks or clay particles) in a volume. Therefore, the volume can be totally occupied by corresponding V-objects (soil blocks or clay

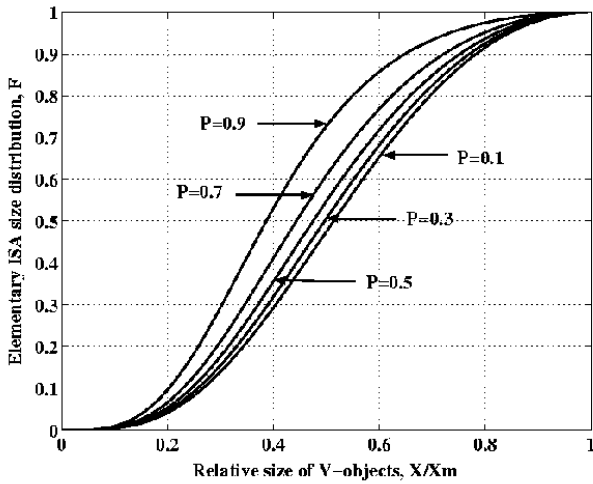


Fig. 4. The numerical presentation of the elementary ISA distribution at different P values.

matrix pores), *ie* the case is possible when the maximum volume fraction of the V-objects, $P_{max}=1$. However, a soil can contain more than one type of S-objects (see ‘Peculiarities...’) and more than one type of V-objects (*eg* pores, sand grains, aggregates). In this case, the volume fraction of V-objects of every type (P_p – for pores; P_s – for sand grains; P_a – for aggregates) should be less than unity. Then, ISA model can be applied to volume parts that are only divided by S-objects of a given type and are only occupied by corresponding V-objects. As a result, we come to the same elementary ISA distribution (Eqs (9) and (10)) for every type of V-objects with replacements $P \rightarrow P_p$, $P \rightarrow P_s$, $P \rightarrow P_a$, and $x_m \rightarrow x_{mp}$, $x_m \rightarrow x_{ms}$, and $x_m \rightarrow x_{ma}$ for pores, sand grains, and aggregates, respectively. With that the following obvious condition should be fulfilled:

$$P_p + P_s + P_a = 1. \tag{12}$$

GENERALIZATION: THE CASE OF SEVERAL MODES OF S-OBJECTS AND CORRESPONDING V-OBJECTS OF A SIMILAR TYPE

Note that the elementary distribution (Eqs (9) and (10)) relates to a certain pair of x_m and P values. However, V-objects of a given type (pores, or sand grains, or aggregates) can be a component mixture with several different pairs of x_m and P values. Such components of the mixture we will refer to as modes. The volume fraction of the modes in such a mixture meet the obvious condition of:

$$P = \sum_{i=1}^I P_i \tag{13}$$

where: I is the mode number of the V-objects of a given type; P_i is the volume fraction of the i -th mode (of the total soil volume). In the case of pores $P \equiv P_p$, $P_i \equiv P_{pi}$, and $I \equiv I_p$. In the case of sand grains $P \equiv P_s$, $P_i \equiv P_{si}$, and $I \equiv I_s$. In the case of aggregates $P \equiv P_a$, $P_i \equiv P_{ai}$ and $I \equiv I_a$. The size distribution of the mixture of V-objects of a given type (pores, or sand grains, or aggregates) can be presented as the sum of I terms of corresponding elementary ISA distributions (Eqs (9) and (10)) multiplied by the weights

$$P_i / \sum_{i=1}^I P_i:$$

$$F(x) = \sum_{i=1}^I [1 - (1 - P_i)^{I_0(x/x_{mi})/8.4}] / \sum_{i=1}^I P_i, \quad 0 < x \leq x_{mI} \tag{14}$$

(in the case of $I=1$ we return to Eq. (9)). We accept that $0 < x_{m1} < x_{m2} < \dots < x_{mI}$. The i -th term of Eq. (14) (in square brackets) is equal to P_i in the range $x_{mi} \leq x \leq x_{mI}$. If the $F(x)$ function gives a pore size distribution, $\sum_{i=1}^I P_i = P$,

where P is the total porosity (cf. Eq. (13)), and P_i is the volume fraction of pores of the i -th mode (of the soil volume). If the $F(x)$ function gives the solids-size distribution (sand grains or aggregates), $\sum_{i=1}^I P_i = 1 - P$,

where: P again is the total porosity, and P_i is a volume fraction of solids of i -th mode (of the soil volume).

One should emphasize that $F(x)$ distribution (Eq. (14)), as the elementary ISA distribution (Eqs (9) and (10)), is a universal function of the ratios x/x_{mi} ($i = 1, \dots, I$) for V-objects meeting the geometrical conditions of ISA. Materials and methods can only influence values of the P_i ($i=1, \dots, I$) and x_{m1}, \dots, x_{mI} parameters.

DATA TO BE USED

Materials and methods

There are many data on the size distributions of pores, sand grains, and aggregates. To illustrate and test the generalized ISA model as applied to pores, sand grains, and aggregates as V-objects, we considered three examples of available data sets relating to sand skeleton pores (Day and Luthin, 1956), sand grains (Mualem, 1976), and aggregates (Wittmuss and Mazurak, 1958). These data are presented as points in Figs 5-7.

The cumulative pore size distribution of Oso Flaco fine sand (the points in Fig. 5) was calculated (Brutsaert, 1966) based on the data for sand-water retention curve (Day and Luthin, 1956) (one of two methods, noted in the Introduction, for finding data on pore size distribution).

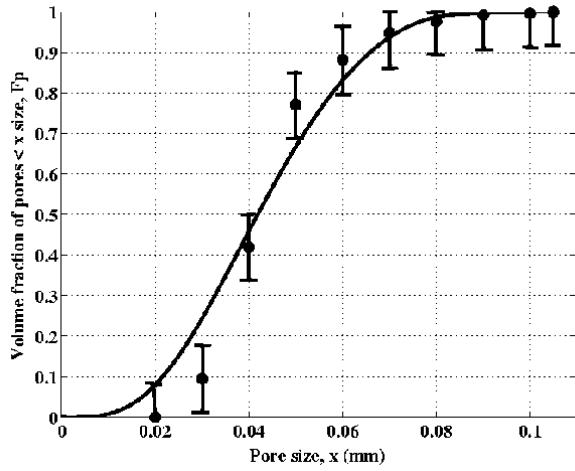


Fig. 5. Data (10 points) of the pore-size distribution of Oso Flaco fine sand from Day and Luthin (1956) (quoted by Brutsaert (1966)) and the fitted distribution curve (solid line) for two elementary pore modes (Eq. (14) at $I=2$): $x_{m1}=0.084$ mm, $P_1=0.375$ and $x_{m2}=0.105$ mm, $P_2=0.03$ (at actual porosity value $P=P_1+P_2=0.405$ from Day and Luthin (1956)). Goodness of fit is $r^2=0.96$. Estimated standard deviation of F_p (vertical bars) is $\sigma=0.08$.

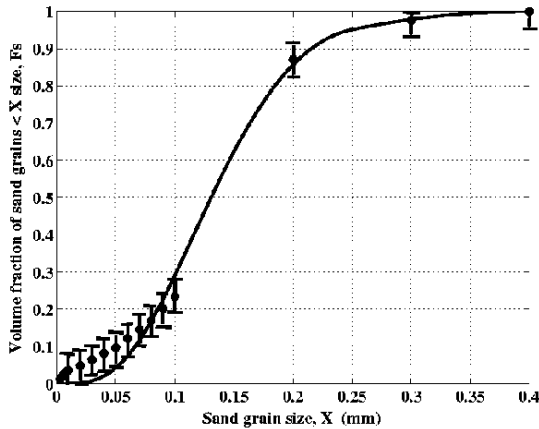


Fig. 6. Data (20 points) of the silty fine sand-size distribution from Mualem's (1976) Fig. 2a and the fitted distribution curve (solid line) for two elementary sand grain modes (Eq. (14) at $I=2$): $x_{m1}=0.25$ mm, $P_1=0.37$ and $x_{m2}=0.4$ mm, $P_2=0.07$ (at actual $x_{m2}=0.4$ mm). Predicted sand porosity is $P=1-P_1-P_2=0.56$. Goodness of fit is $r^2=0.99$. Estimated standard deviation of F_s (vertical bars) is $\sigma=0.04$.

The cumulative sand size distribution from Mualem's (1976) Fig. 2a (the points in Fig. 6) corresponds to the averaged data of silty fine sand obtained by sieving.

Finally, the cumulative aggregate size distribution (the points in Fig. 7) of Sharpsburg silty clay loam (a member of the Brunizem great soil group, from the Agronomy Farm, Lincoln, Nebraska) was determined by sieves, elutriation, sedimentation, and centrifugation (Wittmuss and Mazurak, 1958).

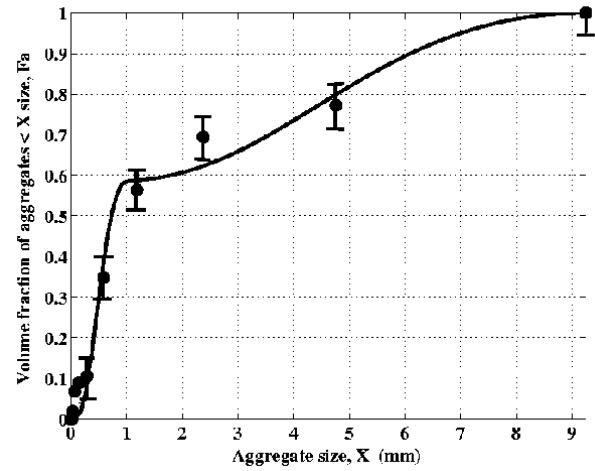


Fig. 7. Data (11 points) of the aggregate-size distribution of Sharpsburg silty clay loam from Wittmuss and Mazurak (1958) (quoted by Rieu and Sposito (1991b)) and the fitted distribution curve (solid line) for two elementary aggregate modes (Eq. (14) at $I=2$): $x_{m1}=1$ mm, $P_1=0.28$ and $x_{m2}=9.25$ mm, $P_2=0.2$ (at actual $x_{m2}=9.25$). Predicted interaggregate porosity is $P=1-P_1-P_2=0.52$. Goodness of fit is $r^2=0.99$. Estimated standard deviation of F_a (vertical bars) is $\sigma=0.05$.

Data analysis

The objective of the data analysis is to show that the above indicated size distributions (the points in Fig.5-7) as functions of size x , can be presented by the $F(x)$ dependence that is characteristic for an ISA distribution (Eq. (14)) despite the differences in the V-object nature, scale, dynamic properties, and preparation methods.

In all three cases (sand skeleton pores, sand grains, and aggregates) we considered the simplest case of two modes, *ie* we approximated the experimental size distributions (10, 20, and 11 points in Figs 5-7, respectively) by the weighted mixtures of two modes. That is, $F(x)$ from Eq. (14) transforms (at $I=2$ and $x_{m1} < x_{m2}$) to:

$$F(x) \equiv F(x/x_{m1}, x/x_{m2}, P_1, P_2) \quad 0 < x \leq x_{m2}. \quad (15)$$

The model was checked by the least squares criterion providing the minimum of the sum of squares, S of the differences between calculated ($F(x_k/x_{m1}, x_k/x_{m2}, P_1, P_2)$) and measured (F_k , see Figs 5-7) values of the size distribution across a range of $k=1, \dots, J$ values of x ($J=10, 20$, and 11 in Figs 5-7, respectively):

$$S \equiv \sum_{k=1}^J [F(x_k/x_{m1}, x_k/x_{m2}, P_1, P_2) - F_k]^2. \quad (16)$$

Analysis of the experimental size distributions (both of pores and solids) using the presentation of Eq. (14) at $I=2$ permits one to estimate P_1 and P_2 (and the total porosity

$P = P_1 + P_2$ or $P = 1 - P_1 - P_2$) as well as x_{m1} and x_{m2} . If P and/or x_{m2} are known in advance (the usual situation), this simplifies the analysis, especially in the cases of $I=1$ or $I=2$. If the standard deviations, σ_k for a number of F_k values are similar and may be considered as a constant, σ the method of the least squares allows one to obtain an independent estimate of σ (Hamilton, 1964) (in addition to the estimates of x_{m1}, x_{m2}, P_1 , and P_2) as:

$$\sigma \cong (S_{\min} / (J - 3))^{1/2} \quad (17)$$

where S_{\min} is the minimum of S as a function of three independent parameters of x_{m1}, x_{m2}, P_1 , and P_2 .

RESULTS AND DISCUSSION

Figures 5-7 show fitted size distributions and estimated σ values (vertical bars). Found values of x_{m1}, x_{m2}, P_1, P_2 and P as well as goodness of fit, r^2 are also indicated in the captions of Figs 5-7. All three examples presented in Figs 5-7 are characterized by high r^2 values. For all three examples, the maximum discrepancies between the experimental points and fitted curve are within the limits of a little more than one standard deviation (σ). These results demonstrate the good agreement between the generalized ISA model and available data despite the differences in the V-object nature (sand pores, sand grains, and aggregates), scale (hundredths of mm in Fig. 5; tenths of mm in Fig. 6; and millimeters in Fig. 7), dynamic properties, and preparation methods.

It is worth emphasizing that the elementary ISA distribution (Eqs (9) and (10)), as well as the generalized distribution (Eq. (14)), is not an empirical one like lognormal or similar distributions. The ISA distributions flow out of the statistics of intersecting surfaces (see Brief summary of ISA). For simplicity we used fitting parameters in the above testing (with the x_{m1}, x_{m2}, P_1 , and P_2 parameters). In general, however, P and x_m parameters (Eqs (9) and (10)), and in the generalized case P_1, \dots, P_I and x_{m1}, x_{mI} parameters (Eq. (14)), have a clear physical meaning and can be measured independently of the data on size distribution (unlike parameters of lognormal and similar empirical distributions). Independent measurement of the P_1, \dots, P_I and x_{m1}, x_{mI} parameters (unlike fitted parameters of empirical distributions) enables one to estimate the size distribution for these parameter values without expensive testing.

According to the ISA approach (see Brief summary), the presentation of Eq. (14) is an analytical basis for the quantitative description of the shape distribution of pores or solids (Eqs (7) and (8)) and the tortuosity of pore channels as a function of P_1, \dots, P_I and x_{m1}, x_{mI} (Eq. (6)). In turn, (pore) size distribution (Eq. (14)), (pore) shape distribution, and tortuosity are necessary for physically based modelling of

the soil water retention curve and hydraulic conductivity function (as an example of such development as applied to the water retention of a clay matrix see (Chertkov, 2004)).

CONCLUSIONS

1. An available approach, connecting the networks of such surface-like objects (S-objects) as cracks and clay particles in soils to the size distribution of corresponding volume-like objects (V-objects) such as blocks and clay matrix pores (the intersecting surfaces approach - ISA), is generalized to the case of different pores, sand grains, and aggregates as V-objects and their boundaries (along with cracks and clay particles) as S-objects. The generalization is based on accounting for the simultaneous presence in a volume of a number of types of S- and V-objects and a number of modes of S- and V-objects of a similar type.

2. For model illustrating and testing we used three examples of available data on the size distributions of sand skeleton pores, sand grains, and aggregates. The model was checked by the least squares fitting of the theoretical distribution curve to the experimental data (in general, the distribution parameters can be found independently). Comparison between the model and available data (based on high values of the goodness of fit and small values of the standard deviations) speaks in favour of the generalized approach.

3. According to the intersecting surfaces approach, the size distributions obtained permit one to construct the shape distributions of pores, sand grains, and aggregates, and calculate the tortuosity of pore channels. All these functions (the size and shape distributions, and tortuosity) are necessary for physically based modelling of soil hydraulic properties.

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